

# Assessment of the effects of the structure on the compression behaviour of a young alluvial silty soil

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## Abstract

A comprehensive series of tests was performed to investigate the effects of the in-situ structure and sample preparation technique on the compression behaviour of a shallow young alluvial silty soil. Two different particle size distributions were investigated by testing intact, compacted, and slurry samples in an oedometer. The compression of the slurry samples, created at different initial water contents, gave a unique intrinsic compression line for one of the two distributions. This was used as a reference for analysing the effects of the structure of the compacted and intact specimens. The compacted samples were prepared by means of static and dynamic compaction in one or more layers with vertical holes simulating those created by plant roots in natural samples. The results showed that a unique normal compression line is defined regardless of the number of layers or the number of holes, with a clear effect of the structure when compared to the slurry samples. The significant effects of the in-situ structure of the intact samples were observed even at high stress levels. An important finding was the non-uniqueness of the compression line for the second distribution. For the case of the slurry samples, the compression lines remained parallel to each other even at high stress levels, whereas a unique line was found for the compacted samples. These results are contrary to those often given in the literature, namely, that the sample preparation technique can create very robust initial structures resulting in a mode of behaviour that has been called “transitional”. The research also emphasises how the quantification of the natural structure is critically dependent on the way in which the intrinsic samples are prepared.

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**Keywords:** Laboratory tests; Structure of young silty soils; Compression behaviour; Quantification of natural structure; Transitional behaviour

## 1. Introduction

Large areas around the world are covered by young alluvial materials. The development and expansion of urban areas on these shallow heterogeneous materials require a rigorous characterization of their geotechnical properties as reliable support for engineers in designing new structures. The effects of the naturally occurring in-

situ structure on their mechanical behaviour are of paramount importance.

The term “structure” in soils is defined as a combination of the particle arrangement, called the “fabric”, and the inter-particle forces, called “bonding” (Mitchell, 1976). Bonding refers to all the inter-particle forces that are not of a purely frictional nature (Cotecchia and Chandler, 1997). In large parts of scientific literature, it is usual to find terms such as “destruction” or “removal of structure” to address the processes that occur in soils when they are strained during testing. This terminology was introduced in soil modelling to account for the effects of the

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structure evolution that take place in natural soils when compared to the structure of reconstituted soils. However, this could lead to the misunderstanding that a soil has “more” structure than the same soil at a reconstituted state which is often said to be “structureless”. In reality, all soils have structure and that structure evolves to a different one during straining. A large part of the geotechnical community uses this terminology to explain the real physical behaviour of soils; and thus, confusion arises. In this research, efforts have been made to avoid this terminology. It is also common to find the term “remoulded” in the literature as being equivalent to that of reconstituted, but they represent different structures, resulting from the sample preparation procedure followed in each case. The term “remoulded soil” defines a state at which the initial natural structure is broken down at the current water content to produce a more homogeneous structure. Different structures are created depending on the energy applied during the remoulding process, which might influence the soil behaviour, e.g., in the case of natural materials with a complex structure (e.g., Dumbleton, 1967; Fearon and Coop, 2000; Madhusudham and Baudet, 2014). Laboratory remoulded compacted samples are also used as a reference material for predicting the behaviour of silty soils compacted in-situ at the same density and the same water content. This comparison has been proven unreliable due to the different structures created during compaction (e.g., Jommi and Sciotti, 2003).

In fine-grained material, the effects of the structure of natural intact samples on the soil behaviour have been evaluated routinely by comparing the intact soil response with that in its reconstituted state (e.g., Burland, 1990; Leroueil and Vaughan, 1990; Cotecchia and Chandler, 1997, 2000). In this comparison, it is assumed that there exists a unique normal compression line (NCL) and a unique critical state line (CSL) for the reconstituted soil state from which the intrinsic parameters can be obtained. The unique NCL was defined by Burland (1990) as the intrinsic compression line (ICL). In his Rankine lecture, he concluded that if the soil is reconstituted at a water content of between  $w_1$  and  $1.5w_1$ , the ICL is well defined for pressures equal to or greater than 100 kPa. The key point of this comparison is whether the NCL and the CSL are unique or not for a given reconstituted soil state.

The non-uniqueness of the NCL and the CSL has been found for a large range of materials with very different gradings and mineralogy for reconstituted and compacted samples (e.g., Martins et al., 2001; Nocilla et al., 2006; Altuhafi and Coop, 2011; Shipton and Coop, 2012, 2015; Xiao et al., 2016; Xu and Coop, 2017). The term “transitional” has been used to refer to this mode of behaviour in reconstituted/remoulded materials where the initial structure of the soil dominates its compression and shearing response even at very high stresses and large strains. Transitional soils represent a clear example of the effect of the robustness of the initial structure, particularly the fabric, during compression and shearing, for which the

term “intrinsic” properties, defined by the reconstituted samples, lacks any practical meaning. Many authors concluded that the sample preparation method alone does not produce different structures that result in a non-convergence of the compression paths; and therefore, they pointed out that the sample preparation method has no effect on the transitional soil behaviour (e.g., Martins et al., 2001; Nocilla et al., 2006; Ferreira and Bica, 2006; Shipton and Coop, 2015). This latter conclusion is investigated in this research as will be discussed below.

Soils of similar gradings and mineralogy retrieved from locations close to those tested in this research displayed transitional behaviour (Nocilla et al., 2006; Nocilla and Coop, 2008). A priori, this finding makes them a potential transitional soil that could invalidate the possibility of evaluating its natural in-situ structure in the traditional way. The materials tested in this research consist of shallow young alluvial silty soils from the plains of the Bormida River at Castellazzo, Italy. These materials are referred to as the Bormida River silts (BRS). Similar data for such young natural alluvial sediments are sparse in the literature; and thus, the effects of the structure on such young sediments have not been well established. Coop and Cotecchia (1995) carried out a series of triaxial tests on intact samples of a relatively young alluvial silty clay, Sibari clay, and concluded that the non-convergence of the compression lines of the intact soil towards the ICL was attributed to a very stable structure even at high stresses. They also noted that the offset of the intact compression lines with respect to that of the ICL increased with the degree of layering in the samples. Rampello and Callisto (1998) found similar behaviour during compression in young alluvial clayey silt sediments from Pisa. Although it is very difficult to determine the different mechanical roles of each component of the structure of a soil, that is, fabric and bonding, Coop et al. (1995) suggested that the fabric provides the stable component of the structure, whereas bonding represents that part which evolves and degrades with strain.

The aim of this research was to give answers to the following questions: (a) Is it possible to find a unique intrinsic NCL for the BRS? (b) How does the sample preparation technique influence the mechanical behaviour of the BRS? (c) How does the in-situ structure influence the behaviour of the BRS? In an attempt to answer these questions, a specific experimental programme was defined based on a series of oedometer tests carried out on intact, compacted, and slurry samples.

## 2. Materials tested and experimental procedures

### 2.1. Materials

The materials tested here consisted of recently deposited alluvial silty soils obtained from the foundation level of the Bormida River embankments at Castellazzo Bormida, Italy. Two types of samples were retrieved, namely, an

intact block sample (B) taken from a depth of 1.4 m below ground level and a remoulded excavated sample (E) taken from nearby. Six samples of the intact material (B) from different parts of the block and two samples of the remoulded one were subjected to sieving and sedimentation tests following the BS standards. The material contained in the block (B) was very homogeneous in terms of size distribution, which gave an almost unique grading curve to all the samples tested. This homogeneity was also found in the remoulded (E) specimens. Fig. 1 shows the average particle size distribution for each material. According to the index properties shown in Table 1, the soils were classified as an inorganic low plasticity clayey silt (B) and an inorganic low plasticity sandy silt (E). A series of X-ray diffraction tests was performed on whole samples and on samples containing only particles smaller than 4  $\mu\text{m}$ . The mineralogy of the silt and sand fractions was mostly quartzitic in nature and the clay fractions were illite and smectite. A visual inspection of the clayey silt intact samples at a macro-structure level showed that organic matter could be found randomly in the form of small plant roots, as shown in Fig. 2. In some cases, vertical holes from plant roots were found that crossed the entire sample. This in-situ structure was reproduced in the laboratory by creating vertical holes in the compacted samples, as will be explained below. The average organic matter content was only 1.1%; and thus, its effects on the mechanical behaviour of the clayey silt (B) are likely to have been insignificant



Fig. 2. Plant root in an intact sample of clayey silt (B).

(e.g., Booth and Dahl, 1986). Negligible quantities of organic matter were measured in the sandy silt (E) soil.

## 2.2. Sample preparation technique

Three types of oedometer samples were used to investigate the influence of the structure and sample preparation method on the mechanical behaviour of the BRS soils.

### • Intact samples

The block sample of the clayey silt (B) was divided into small pieces that were large enough to use as samples for the oedometer tests. The position of each piece within the block was identified to check the homogeneity of the material contained in the block since, for alluvial soils, rapid changes in the depositional environment can make them highly heterogeneous (e.g., Coop and Cotecchia, 1997). Intact specimens were trimmed from each piece of material. The oedometer ring, with a sharp edge, was placed on top of the soil piece and small pressure was applied on the ring, pushing it against the sample at the same time as the trimming was carried out, trying to advance the ring and cutting at the same time to avoid disturbance. Once the sample was inside the ring, the top and bottom were carefully trimmed, paying special attention to create a very flat surface on both sides. During the trimming process, all the remaining material was collected and later used for

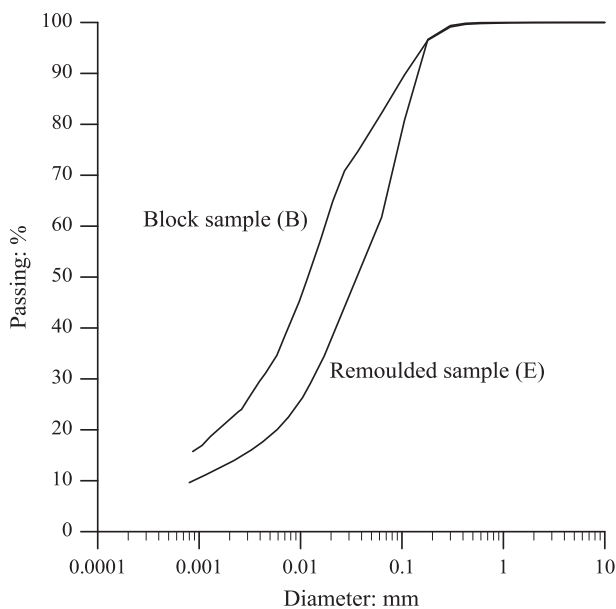


Fig. 1. Particle size distributions of Bormida River silts (BRS).

Table 1  
Index properties.

Soil	Clay-size: %	Silt: %	Sand: %	LL	PI	USCS	A	OM: %	CaCO <sub>3</sub> : %
Intact block (B)	22	60	18	35.2	10.3	ML	0.47	1.1	10.4
Remoulded (E)	13	50	37	25.3	4.2	CL-ML	0.33	–	12

Table 2  
Summary of oedometer tests carried out on samples of clayey silt (B) and sandy silt (E).

Test	Preparation method	Sample diameter (mm)	Initial void ratio $e_0$	Yield stress $\sigma'_y$ : kPa	$\sigma_v$ : kPa <sup>a</sup>	Comments
B_S1	Slurry	50	1.050	–	–	
B_S2	Slurry	50	1.105	–	–	
B_S3	Slurry	38	0.993	–	–	
B_S4	Slurry	38	1.077	–	–	
B_S5	Slurry	38	1.274	–	–	
B_DC1	DC_W	38	0.590	840	–	
B_DC2	DC_W	50	0.826	130	–	
B_SC3	SC_W	50	1.132	20	86.6	1 layer
B_SC4	SC_W	50	0.849	70	407.5	1 layer
B_SC5	SC_W	50	0.790	160	–	1 layer
B_SC_L1	SC_W	50	0.828	170	–	2 layers; $e_1 \cong 0.6$ , $e_2 \cong 1$
B_SC_L2	SC_W	50	0.853	120	–	3 layers; $e_1 \cong 0.6$ , $e_2 \cong 0.8$ , $e_3 \cong 1$
B_SC_H1	SC_W	50	0.867	145	–	3 drilled holes
B_SC_H2	SC_W	50	0.828	160	–	3 pushed holes
B_SC_H3	SC_W	50	0.878	120	407.5	17 drilled holes
B_SC_H4	SC_W	50	1.102	30	204	10 drilled holes
B_SC_H5	SC_W	50	1.140	24	86.6	14 pushed holes
B_I1	Intact	50	0.808	300	–	
B_I2	Intact	50	0.765	300	–	
B_I4	Intact	50	0.713	345	–	
B_I5	Intact	50	0.748	230	–	
B_I6	Intact	50	0.778	365	–	
B_I7	Intact	50	0.721	300	–	
B_I8	Intact	38	0.723	400	–	
E_S1	Slurry	50	0.807	–	–	
E_S2	Slurry	38	1.031	–	–	
E_S3	Slurry	38	0.703	–	–	
E_DC1	DC_W	50	0.635	900	–	
E_DC2	DC_A	50	0.793	250	–	

Note: B: clayey silt; E: sandy silt; S: slurry; I: intact; DC: dynamic compaction; SC: static compaction; DC\_W: dynamic wet compaction; SC\_W: static wet compaction; DC\_A: dynamic air-dry compaction.

<sup>a</sup> Maximum vertical stress applied during static compaction.

preparing the slurry and remoulded samples. The intact samples are identified as B\_I in Table 2.

- Slurry samples

Slurry specimens were prepared by adding distilled water to the trimmings from the intact samples of the clayey silt (B) and the remoulded samples of the sandy silt (E) and mixing thoroughly to achieve a uniform consistency. Following Burland (1990), the soil trimmings were not oven-dried before mixing. Different specimens were prepared by adding different amounts of water from the liquid limit to 1.5 times the liquid limit. Special attention was paid to the possible segregation of the sandy silt (E) during the mixing process. The slurry was placed inside the oedometer ring avoiding any air entrapment. The slurry samples tested in this research are identified in Table 2 as B\_S and E\_S for the block and the remoulded excavated materials, respectively.

- Compacted samples

In the case of the clayey silt (B), the trimmings from the intact specimens were remoulded by hand using a spatula,

maintaining the initial water content of the intact samples. The same procedure was followed for the sandy silt (E). The samples were compacted inside the oedometer ring either statically or dynamically. The dynamically compacted (DC) samples were prepared in three layers. In the case of the statically compacted (SC) specimens, different samples were created, such as samples compacted in one layer, samples compacted in one layer with vertical drilled holes, samples compacted in one layer with vertical pushed holes, and samples compacted in several layers but with a very different value for the void ratio of each layer. These latter samples were called “layered samples”. Compaction of the statically compacted (SC) specimens was conducted using specifically designed equipment, as shown in Fig. 3 (a). Each sample was compacted inside the ring by applying a constant rate of displacement to the base pedestal. A static flat-ended rod, connected to a load cell, was used to create the reaction to the application of static pressure to the sample and the maximum static load was recorded. For the layered samples, the surface between each layer was scarified before continuing with the next layer to ensure good continuity in the whole sample. All the samples were allowed to rest for 5 min before placing the oedometer ring in the cell in order to permit them to rebound. Both the



(a)



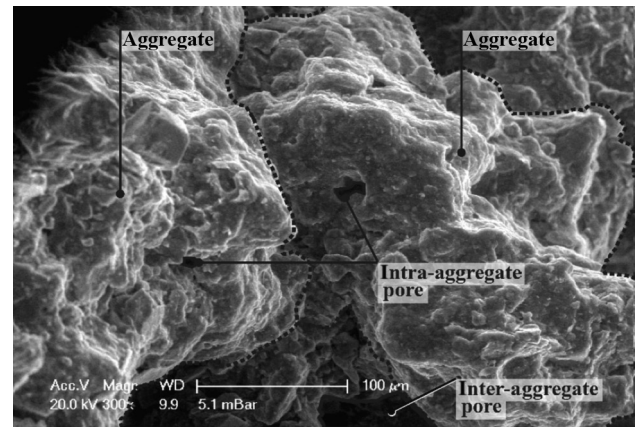
(b)

Fig. 3. (a) Equipment used for static compaction and (b) oedometer compacted samples with drilled holes in a specimen of clayey silt (B).

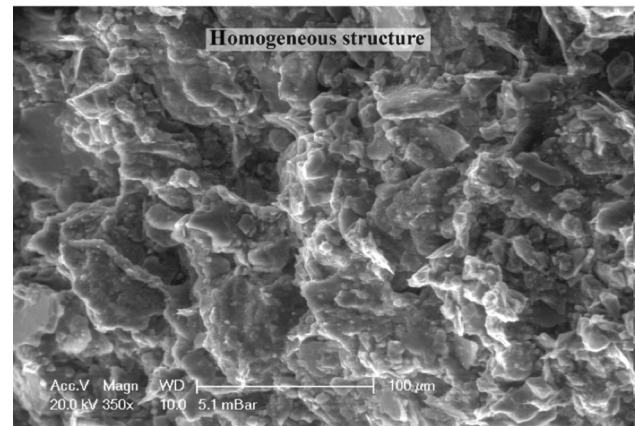
bottom and top of the samples were trimmed to be flat before starting the test.

Holes and plant roots were often found during the trimming of the intact specimens of the clayey silt (B), as described above and shown in Fig. 2. They were mainly vertical holes and, in some cases, they extended from the top to the bottom of the sample. In order to reproduce these vertical holes, drilled or pushed holes were artificially created. The difference between the two procedures is that the drilled holes were created by extracting material from the sample and the pushed ones by displacing the material inside the sample. A temporary top platen was used to avoid any sample volume change during the drilling and pushing procedures. The number and the location of the holes inside the sample were randomly distributed. Fig. 3 (b) shows an oedometer specimen after the holes had been drilled.

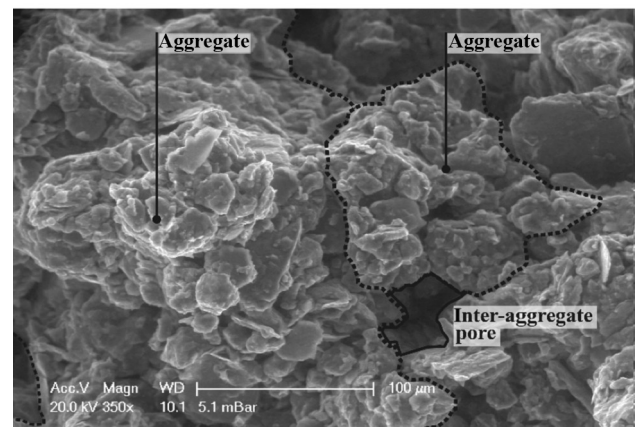
Micro-structure studies were conducted with SEM images taken from different specimens of the clayey silt (B). Fig. 4 shows the initial micro-structures of the intact, slurry, and compacted samples before testing. The micro-structure of the intact sample is characterised by a fabric



(a)



(b)



(c)

Fig. 4. Scanning electron microscopy images of: (a) intact, (b) slurry, and (c) compacted samples of clayey silt (B).

where clear aggregates of particles can be identified; they are bonded by bridges of particles of smaller sizes. This aggregation of soil particles resulted in an open structure with large inter-aggregate pores. In the case of the slurry sample, a more homogeneous structure can be distinguished, so it appears that the aggregates of the particles found in the intact specimen were destroyed during the reconstitution process to create the slurry. The disturbance of the intact structure towards a more homogeneous one seems to be less efficient in the case of the compacted specimen. In this case, the intact material was remoulded at its natural water content, mixing by hand with a spatula. The structure of the remoulded sample exhibits a similar aggregated structure to that of the intact sample, with a bi-modal pore size distribution characterized by large inter-aggregate pores and smaller intra-aggregate pores.

### 2.3. Experimental procedures

The oedometer tests were performed by applying incremental loading, using oedometer rings with a diameter of 50 mm and a height of 20 mm. In some cases, a 38-mm ring was used to reach higher stress levels in the range of 12–13 MPa compared with the 8–9 MPa for the 50-mm ring. For both the intact and the compacted specimens, an initial load that ranged from 2 to 52 kPa was applied; and thereafter, the cell was filled with water, varying the initial load in order to study its effect on the swelling properties. After filling the cell with water, the samples were left for 24 h before applying the next load. During this time, the swelling behaviour was monitored. It was common to see that the volume change due to swelling stabilised very quickly, although, as mentioned above, the samples were all subsequently left for 24 h to allow them to saturate.

In the case of the slurry samples, after the preparation of each specimen, the soil was placed directly into the oedometer ring paying special attention so as not to trap any air. After applying the first load, the water bath was flooded. Some of the oedometer cells were modified in order to be able to flush water through the sample, at a pressure of 10 kPa. This procedure was applied to some intact and compacted specimens to compare the results with the conventional procedure where the water bath was only flooded to saturate the sample, but no differences were observed. Special attention was paid when dismantling the cell at the end of the test, as the water in the cell could be sucked into the sample after the last load was released. To avoid this, all the water that surrounded the sample and porous stones was dried before dismantling the cell to minimise the error in the final water content.

The calculation of the initial void ratio ( $e_o$ ) of a sample is most important when analysing the results of an oedometer test. Accurate  $e_o$  are essential when investigating the possible effects of the structure to ensure the correct location of the compression data. Eqs. (1)–(4) were derived to calculate the initial void ratio of the sample at the start of the oedometer test and before the first load was applied, where  $G_s$  is the

specific gravity,  $w_i$  and  $w_f$  are the initial and final water contents, respectively,  $\varepsilon_v$  is the volumetric deformation measured during the whole test,  $\rho_{di}$  is the initial dry density, and  $\rho_{bi}$  is the initial bulk density. They are related to the independent parameters that can be measured (Rocchi and Coop, 2014). These independent measurements are the initial water content, the initial wet mass, the initial height, the initial diameter, the final water content, and the final dry mass. Eqs. (2) and (3) were only applied for the slurries, where the samples were assumed to be fully saturated at the start of the test. When the sample was saturated from the start, as in the case of the slurries, all four equations were used and an average value was obtained. For the intact and compacted samples, as they were not initially saturated, only Eqs. (1) and (4) were applied.

$$e_o = \frac{G_s \times \rho_w}{\rho_{di}} - 1 \quad (1)$$

$$e_o = G_s \times w_i \quad (2)$$

$$e_o = \frac{\rho_s - \rho_{bi}}{\rho_{bi} - \rho_w} \quad (3)$$

$$e_o = \frac{w_f \times G_s + \varepsilon_v}{1 - \varepsilon_v} \quad (4)$$

## 3. Experimental results

### 3.1. Compression behaviour of the clayey silt (B)

#### • Slurry samples

A series of oedometer tests was performed on the slurry specimens prepared at different initial water contents from around 1–1.5 times the liquid limit, using the trimmings. Table 2 shows the details of each test. The objective was to investigate the influence of the initial water content of the slurry samples on the uniqueness of the normal compression line. The compression behaviour of the slurry samples is shown in Fig. 5. Although there is some small scatter in the compression lines, a clear tendency towards a unique normal compression line can be seen, regardless of the initial water content, and therefore, the initial structure of the specimens created during the reconstitution process. The small scatter could be associated with the accuracy in measuring the initial void ratio of the samples. An estimated unique normal compression line is drawn in Fig. 5. Following the definition given by Burland (1990), this is the one-dimensional intrinsic compression line (1D-ICL<sup>\*</sup>) of the clayey silt (B) whose slope is  $C_c^* = 0.25$ . It is important to point out that the estimated straight 1D-ICL<sup>\*</sup> deviates from that of the curved shape of the ICL proposed by Burland. During swelling, the behaviour is essentially the same in all the samples with an average swelling index of  $C_s^* = 0.05$  which gives a ratio of  $C_s^*/C_c^* = 0.2$ . The asterisk denotes the intrinsic properties of the material which are inherent to the soil and are independent of the initial structure (Burland, 1990). These intrinsic

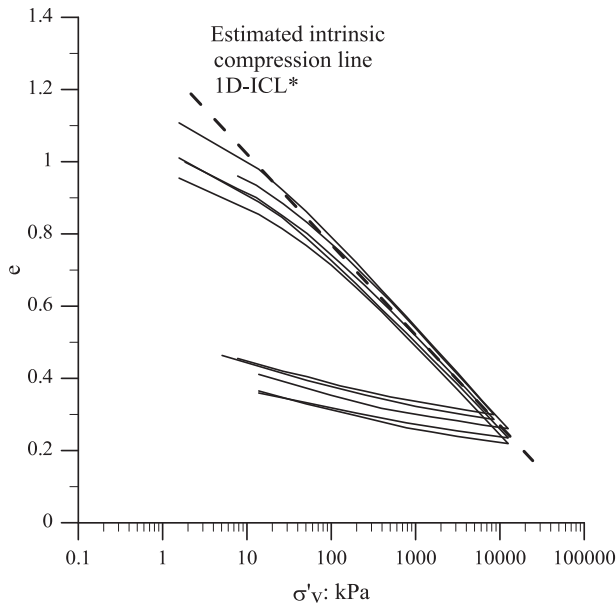


Fig. 5. Oedometer compression lines of slurry samples of clayey silt (B).

properties will be used as a reference for assessing the initial structure of the intact and compacted samples. Burland proposed an equation to correlate  $C_c^*$  with the void ratio at the liquid limit ( $e_{LL}$ ) for soils above the A line on the plasticity chart. The plasticity characteristics of the clayey silt are plotted slightly below and very close to the A line on the plasticity chart and, although the equation proposed by Burland was derived for clays, a good approximation is obtained with a value of  $C_c^* = 0.21$ .

• Compacted samples

Table 2 shows the details of the tests conducted on the compacted samples. The layered samples were compacted to have a different initial void ratio for each layer to create a heterogeneous structure (Table 2). These ranges were to simulate the slightly heterogeneous void ratios within the natural samples, although more extreme variations were used with  $e_o$  from 0.6 to 1.0. It is important to point out that the definition of a layered soil given here is not the same as the one that can usually be found in the literature, where a layered sample consists of soils with different gradings, e.g., layers of clay and sand such as varved clays. The compression lines of the compacted samples shown in Fig. 6(a) converge onto a unique NCL regardless of the compaction method used and the number of soil layers created during compaction. This is surprising given the large differences in the  $e_o$  of each layer, which varied from 0.6 to 1.0 (Table 2). Therefore, although the initial structure of each sample was expected to be different, the deformation induced during compression was enough to ensure that no differences in structure would be sufficient to significantly affect the macro-behaviour of the soil. This result is important in terms of the robustness of the unique NCL regardless of the heterogeneities created at meso-structure and micro-structure levels during the sample preparation

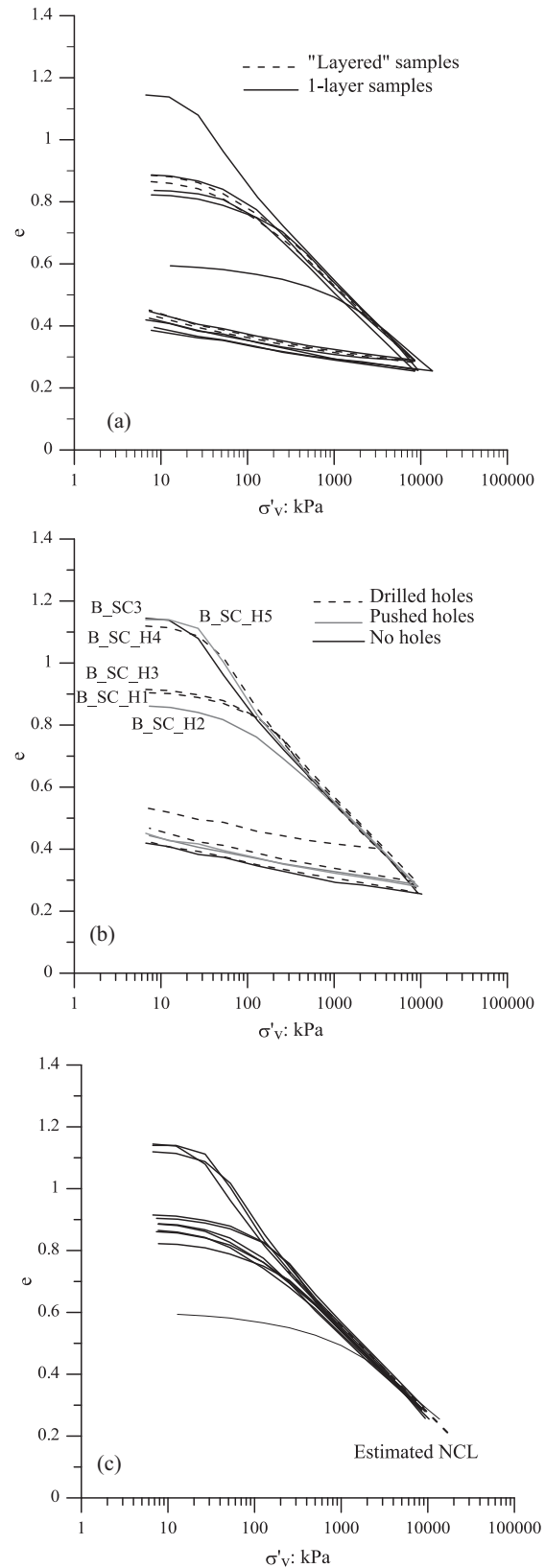


Fig. 6. Oedometer compression lines of compacted samples of clayey silt (B): (a) layered, (b) holes, and (c) all samples.

process. Moreover, this does not mean that the structure of each sample became the same at the end of the test.

As described above, some of the intact samples tested in the oedometer had vertical holes inside. In order to simulate the possible effect of these holes on the compression behaviour, a few oedometer tests were conducted on statically compacted samples with vertically drilled and pushed holes (Table 2) (Fig. 3b). The vertical holes were randomly distributed across the specimens. As described before, the main difference between the drilled holes and the pushed holes was that the drilled holes were created by extracting material from the sample and the pushed holes were created by displacing the material inside the sample, possibly creating a more dense structure around the holes. A plant root might tend to displace rather than remove soil. Fig. 6(b) shows the compression behaviour for all the samples with holes together with that of a compacted specimen without holes, B\_SC3, which is included for comparison. A clear unique compression line is identified; it shows that the vertical holes had no significant effect on the compression response of the clayey silt at higher stress levels. It is also shown in Fig. 6(b) that the number of holes created for the same procedure had no effect either, as can be observed when comparing the compression lines for samples B\_SC\_H1 (3 drilled holes) and B\_SC\_H3 (17 drilled holes). A comparison of the compression lines of the compacted sample without holes, B\_SC3, and the compacted specimens with holes, B\_SC\_H4 and B\_SC\_H5, shows that for a similar average initial void ratio, there was no significant effect of the holes on the soil behaviour. Consequently, it appears to be most unlikely that the root holes found in some natural samples have any significant impact, by themselves, on their behaviour.

Fig. 6(c) shows the compression lines for all the compacted samples tested. The unloading behaviour has not been included for clarity. As can be observed, the specimens with the highest  $e$  values display a stiff response at the start of the compression stages followed by an abrupt change in their compressibility after yield. The change in stiffness after yield resembles the behaviour of bonded sensitive clays, but the effect of bonding is expected to be negligible here as a consequence of the remoulding process. As the initial density of the specimens increases, this apparent effect of the initial structure is less evident. For stresses beyond yield, all the curves converge towards a unique NCL regardless of the initial structure.

Traditionally, for reconstituted fine-grained materials, the yield stress ( $\sigma'_y$ ) observed during one-dimensional compression has been associated with the maximum stress level to which the soil has been subjected during its history, called the preconsolidation pressure ( $\sigma'_p$ ). In the case of uncemented coarse-grained materials, the yield stress occurs at the onset of major particle breakage. During the preparation process of the statically compacted samples, it was possible to measure the maximum vertical stress applied to some of the specimens and identified as  $\sigma_v$ ; kPa<sup>a</sup> in Table 2. The yield stress for each test was estimated using the Onitsuka (1995) method, the values being compared with the maximum compaction stresses in Table 2.

It can be seen that there is no correlation between the maximum vertical stress applied during compaction and the yield stress, meaning that the soil did not “remember” the maximum stress to which it had been subjected. In fact, the maximum vertical stress applied during compaction cannot be compared with the yield stress because the samples were not saturated during compaction; and therefore, the vertical effective stress applied is not known, as is expected.

An attempt to correlate the initial void ratio of each sample after soaking with the measured yield stress is shown in Fig. 7. Although there is some scatter in the results, a relatively good correlation is found with an  $R^2 = 0.98$ . The slope of the linear regression in the  $e$ : $\log \sigma'_y$  plane is 0.34; this is equivalent to the average value of the slope of the estimated unique NCL of the compacted samples plotted in Fig. 6(c). It could be said that the comparison of the two slopes is another way of trying to check the uniqueness of the NCL and also shows that the yield stress of the compacted samples is simply correlated with the initial density or void ratio.

#### • Intact samples

One of the main objectives of this research was to investigate the effect on the compression behaviour of the in-situ structure of the recently deposited alluvial clayey silt (B). Table 2 summarises the details of each oedometer test performed on the intact samples. Although the intrinsic index properties of the material contained in the block sample were the same for each of the samples tested, its in-situ structure seems to be slightly more heterogeneous as revealed by the differences in the initial void ratios, which range from 0.713 to 0.808 with an average value of 0.751

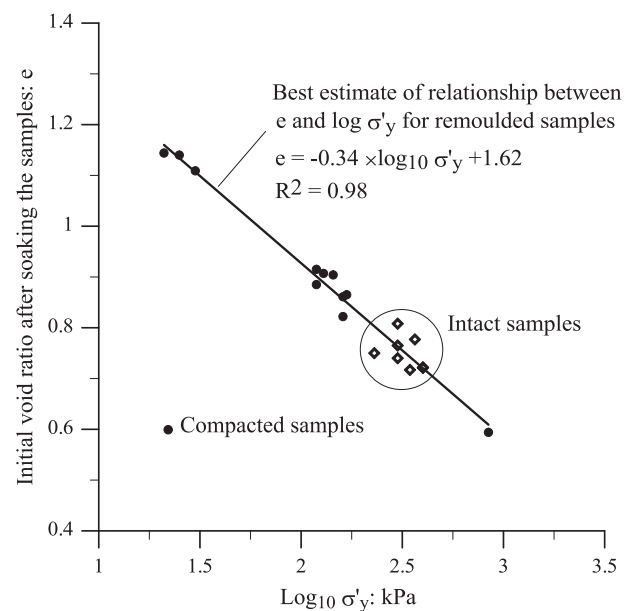


Fig. 7. Correlation between initial void ratio after soaking and yield stress for compacted and intact samples of clayey silt (B).



(Table 2). This variability in the initial void ratio might be associated with a combination of different factors such as a possible heterogeneous initial structure inside the block, the error introduced in calculating  $e_o$ , and the structural damage induced to the samples during the preparation process. These latter two factors are unlikely to be the main source of the variability in  $e_o$  due to the very small error obtained in the calculation of the initial void ratio and the very careful trimming of the samples. During the trimming of the specimens, the granulometry inside the specimen was found to be homogeneous with no layering of materials with different gradings, supporting the idea that the variability of the initial void ratio is not associated with the granulometry, but is most likely to be associated with a possible layered structure. Another important factor that influences the range in the initial void ratio is the presence of the root holes.

The compression behaviour of the intact samples is shown in Fig. 8. The initial parts of the compression lines run almost parallel to each other, with a very stiff response similar to the behaviour of a coarse-grained material, where the initial structure controls the compression response, particularly the fabric. Consequently, in the range of engineering stress levels (<1 MPa), a non-unique compression behaviour can be identified. After the yield, no abrupt changes are observed in the slope of the NCLs, with the samples displaying a slow tendency to converge towards a unique NCL at high stresses, exhibiting a clear effect of the robust initial natural fabric.

Initially, the intact samples were partially saturated with a degree of saturation that varied from 62 to 80%. A suction of 90 kPa was measured in two of the samples using the suction probe technique (Ridley and Burland, 1993). In order to saturate the samples, the oedometer cell was

flooded with water after applying the first load. A different procedure was followed for specimen B\_I8, where it was subjected to a circulation of water with a pressure difference of 10 kPa across the sample for one week after applying the first load and flooding the oedometer cell. A comparison with sample B\_I4 shows that the compression behaviour is the same regardless of the saturation method applied.

The initial swelling response of the intact samples was also measured by monitoring the axial displacement of the specimens after the first load was applied and the oedometer cell was flooded with water. Fig. 9 shows the relationship between the oedometric tangent modulus ( $M_{\text{oed}}$ ) and the vertical effective stress ( $\sigma'_v$ ) of some of the intact samples. It can be observed that all the  $M_{\text{oed}}-\sigma'_v$  curves converge onto a single line at stresses higher than the estimated yield values. The effect of the initial swelling due to flooding is also seen. The initial in-situ structure of the intact samples was weakened during swelling with a consequent decrease in initial stiffness as the swelling increased. This shows that part of the initial structure was disturbed during swelling, but that the effect disappeared quickly even at stresses of about 20 kPa.

The yield stresses for the intact samples were also estimated, the values of which ranged from 230 to 400 kPa, as shown in Table 2. All the yield stresses were much higher than the estimated in-situ vertical effective stress of the block sample of  $\sigma'_v = 75$  kPa. The yield stress ratio (YSR), defined as the ratio between the yield stress and the in-situ vertical effective stress, varied from 3 to 5.3. It is very unlikely that the in-situ sample had a maximum past stress level of the same value as the measured yield stress; and therefore, it is believed that the measured yield stresses cannot be correlated with the pre-consolidation pressure.

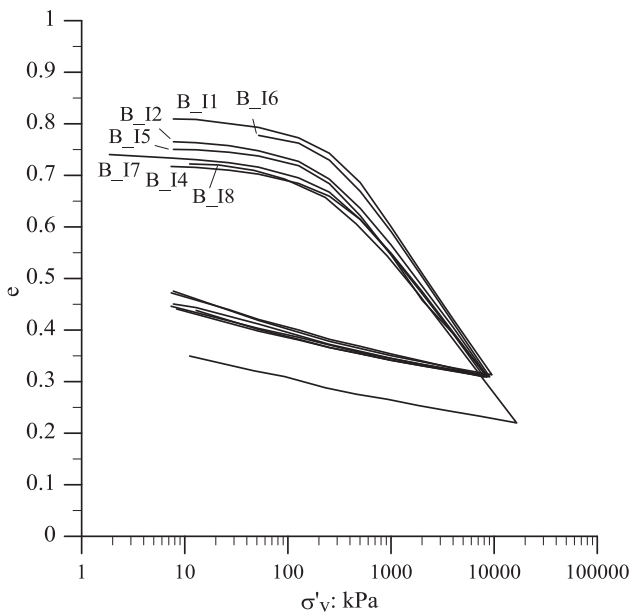


Fig. 8. Oedometer compression lines of intact samples of clayey silt (B).

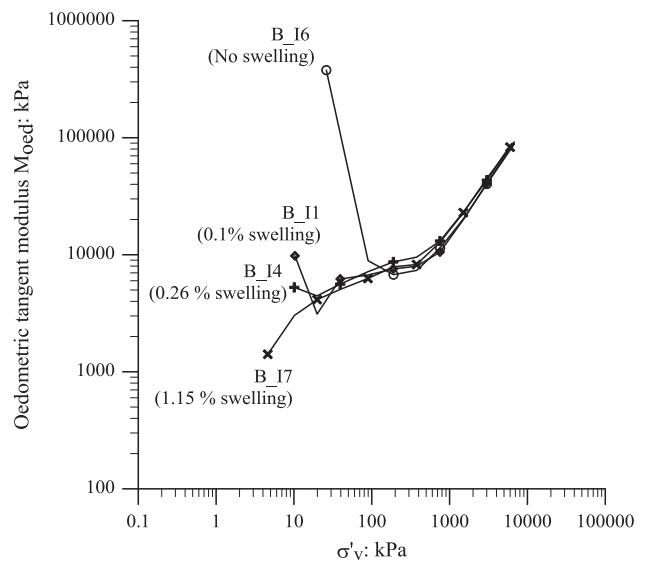


Fig. 9. Oedometer tangent moduli ( $M_{\text{oed}}$ ) of intact samples of clayey silt (B). Effect of initial swelling on compression behaviour.

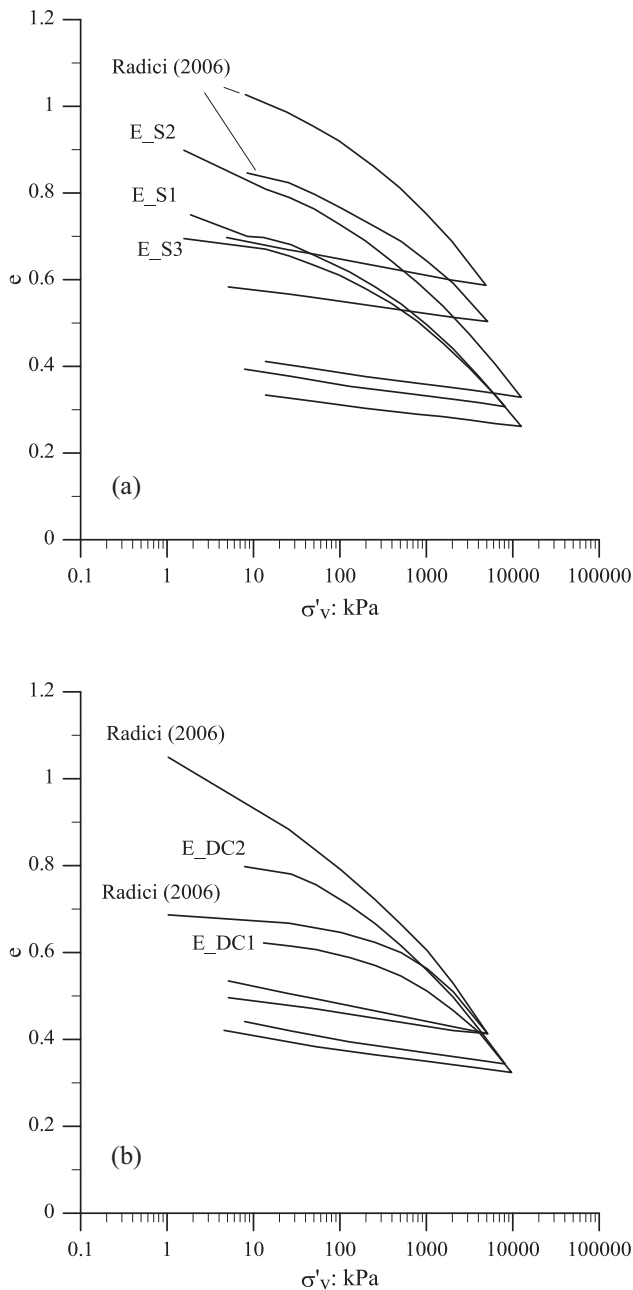


Fig. 10. Oedometer compression lines of: (a) slurry and (b) compacted samples of sandy silt (E).

One factor that could have had an important effect on the high values of the yield stress could have been the desiccation processes due to the shallow depth (1.4 m) from where the block was retrieved. In fact, the liquidity index ( $I_L$ ) values of the intact specimens were negative with an average value of  $I_L = -0.6$ . The yield stresses of the intact samples are plotted in Fig. 7 together with the values measured for the compacted samples. A slightly larger scatter is seen in the results; this scatter could be a consequence of the variability in the in-situ structure, but the data show the same trend and could indicate that the effect of the bonding on the intact samples is small.

### 3.2. Compression behaviour of the sandy silt (E)

#### • Slurry samples

Oedometer tests were carried out on the slurry samples of the sandy silt (E). The slurries were created at different initial water contents, which varied from 1 to 1.5 times the liquid limit. The initial void ratio for each water content is shown in Table 2. The compression behaviour is shown in Fig. 10(a) together with previous results obtained by Radici (2006). Apart from tests E\_S1 and E\_S3, the compression lines do not show any tendency to converge onto a unique NCL despite being tested to stress levels in the range of 8–12 MPa. The position of each compression line depends on the initial void ratio, meaning that the initial structure could not be disturbed enough to reach a unique NCL during compression, even at high stresses, and therefore, showing a clear transitional behaviour. In the case of a relatively small difference in the initial void ratio (e.g., E\_S1 & E\_S3), an apparent convergence could be found. However, this does not mean that for larger differences in initial void ratios there will be the same convergence, as shown with the other specimens. Consequently, it is important to create samples with larger differences in the initial void ratio in order to check the possible transitional behaviour of a soil, but paying attention to avoid any segregation of the soil at water contents that are too high.

Fig. 11 shows a comparison of the compression responses of the slurry samples of the two BRS materials. Only a few compression tests for the clayey silt (B) have been included for clarity. It can be clearly seen that while the compression lines of the clayey silt (B) converge onto a unique 1D-ICL\*, regardless of the initial structure, the response of the sandy silt (E) displays transitional

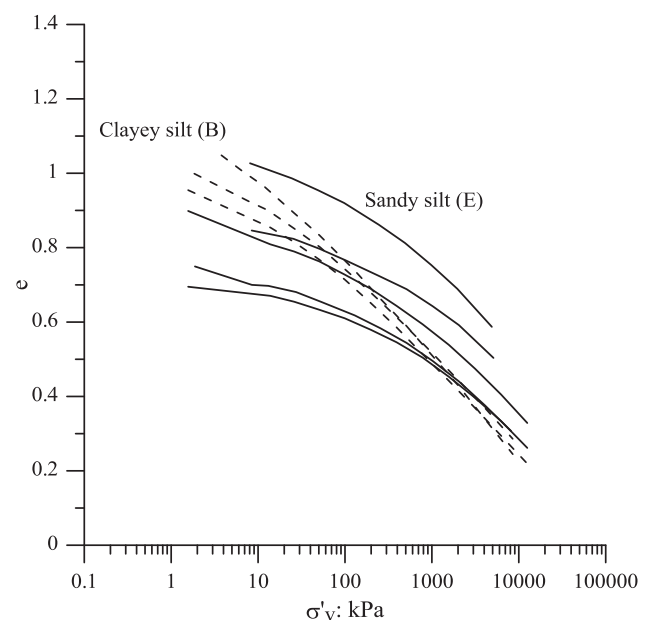


Fig. 11. Comparison of oedometer compression lines of slurry samples of clayey silt (B) and sandy silt (E).

behaviour. The slope of the clayey silt compression lines is slightly higher than that of the sandy silt, possibly due to its higher plasticity.

- Compacted samples

The details of each test conducted on the compacted samples of the sandy silt (E) are shown in Table 2. Fig. 10(b) shows the compression response together with the results of two more tests carried out by Radici (2006). It can be observed that, regardless of the compaction method, all the compression curves converge onto a unique NCL, with a compression index of  $C_c = 0.29$ . These results are in contrast to the behaviour of the slurry samples, for which it is not possible to find a unique NCL during compression for the same range in applied stresses. Consequently, the sample preparation technique must have produced different initial structures, which resulted in a non-convergence of the compression curves, and therefore, determined whether or not there was transitional behaviour of the Bormida sandy silt (E).

### 3.3. Influence of the initial structure on the compression behaviour of the BRS soils

- Clayey silt (B)

In order to evaluate the effects of the in-situ structure on the one-dimensional compression behaviour, the intact soil response is compared with its reconstituted state. In the case of the clayey silt (B), a unique intrinsic compression line, 1D-ICL\*, is found for the slurry specimens and it has been taken as the reference line for comparison with the behaviour of the intact samples. The compression lines of the intact samples and the 1D-ICL\* are normalized by volume using the void index ( $I_v$ ) (Burland, 1990) and are plotted in Fig. 12(a), together with the sedimentary compression line (SCL) defined by Burland. The void index is calculated using Eqs. (5) and (6).

$$I_v = \frac{e - e_{100}^*}{C_c^*} \quad (5)$$

$$C_c^* = e_{100}^* - e_{1000}^* \quad (6)$$

where  $C_c^*$  is the intrinsic compression index determined on the 1D-ICL\*, and  $e_{100}^*$  and  $e_{1000}^*$  are the intrinsic void ratios corresponding to  $\sigma'_v = 100$  kPa and  $\sigma'_v = 1000$  kPa, respectively.

The normalized compression lines of the intact specimens cross the 1D-ICL\* and yield at different stress levels in between the 1D-ICL\* and the SCL. The ratio of the yield stress ( $\sigma'_y$ ) to the equivalent effective stress on the intrinsic compression line ( $\sigma'_{ve}$ ) was defined by Cotecchia and Chandler (2000) as the stress sensitivity ( $S_\sigma = \sigma'_y / \sigma'_{ve}$ ). For the clayey silt (B), the stress sensitivity varies from 1.2 to 2.5 with the lowest and highest values corresponding to the densest and loosest samples, respectively. The post-yield

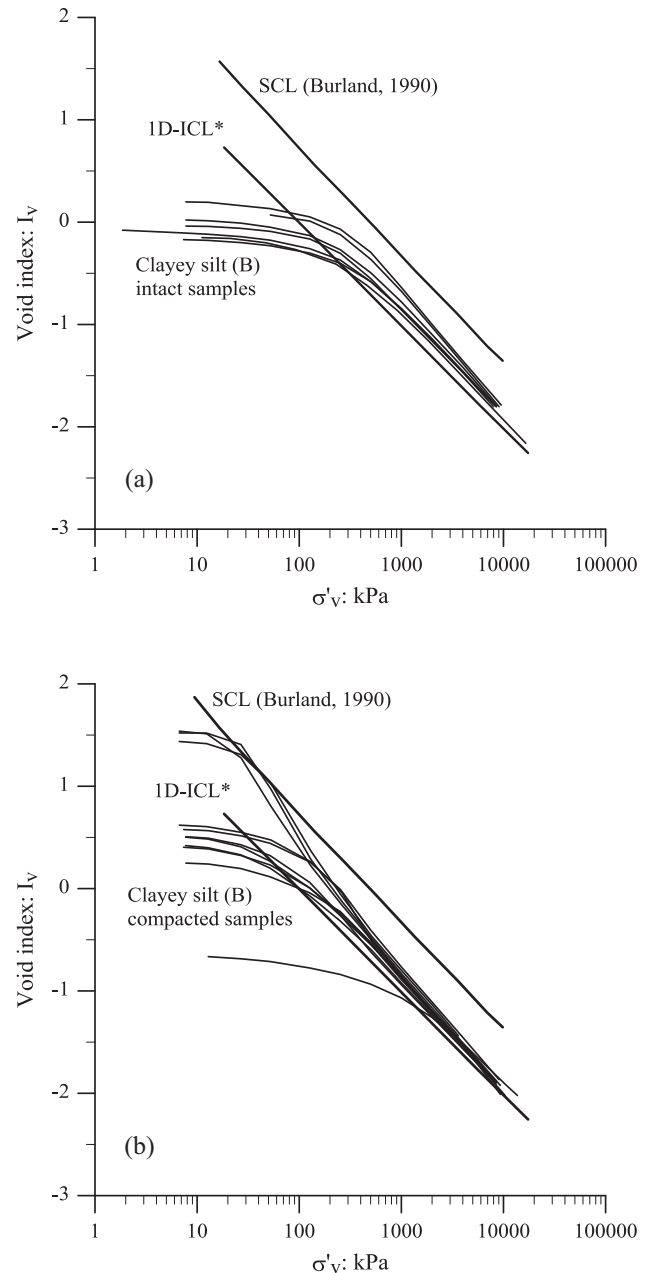


Fig. 12. Comparison between normalized compression lines, measured 1D-ICL\*, and SCL of clayey silt (B): (a) intact samples and (b) compacted samples.

behaviour shows a slow convergence towards the 1D-ICL\*, displaying a robust effect of the initial structure during one-dimensional compression. Burland et al. (1996) suggested that in cases where the yielding stress is located between the 1D-ICL\* and the SCL in a  $\log \sigma'_v : I_v$  plane, the main difference in the compression behaviour between reconstituted and intact samples is due to different fabrics, with a small influence of bonding. It is believed that this is the case of the intact specimens of the Bormida clayey silt (B).

The normalised compression response of the compacted specimens is shown in Fig. 12(b). The compression lines also cross the estimated 1D-ICL\* showing an effect of the initial structure created during compaction. This effect is

more important as the initial density decreases. In fact, the samples with higher initial void ratios reach the SCL with stress sensitivity values of around 5, displaying a post-yield convergence towards the 1D-ICL\*. At high stress levels, the compression lines are plotted only slightly above the 1D-ICL\*. The observed effect of the initial structure of the remoulded compacted samples, when compared with the slurry behaviour, proves that they should not be used as a reference for the analysis of the effects of the initial structure of the natural intact samples. In fact, the intact and compacted effects of the structure are similar and that could lead to the conclusion that there are “no” structure effects at a meso-level if the compacted behaviour is used as a reference. In addition, as in the case of the intact samples, the fabric must be the main part of the structure controlling the soil’s behaviour since there can be no bonding. This result emphasises how the quantification of the natural structure is critically dependent on the way in which the reference samples are prepared.

- Sandy silt (BRS-E)

In the case of the Bormida sandy silt (E), only slurry and compacted samples were tested with no intact specimens available. As shown in Fig. 13, for the slurry samples, it was found that a different NCL is obtained for each initial water content, showing a clear transitional behaviour. The slope of each NCL defined at high stresses is essentially the same. This lack of uniqueness of the NCL would make any comparison between the natural intact samples and the slurry samples to determine the effects of structure impossible, and the concept of intrinsic compression line as a reference framework cannot be used for this more coarse-grained material. On the other hand, a unique NCL is identified for the com-

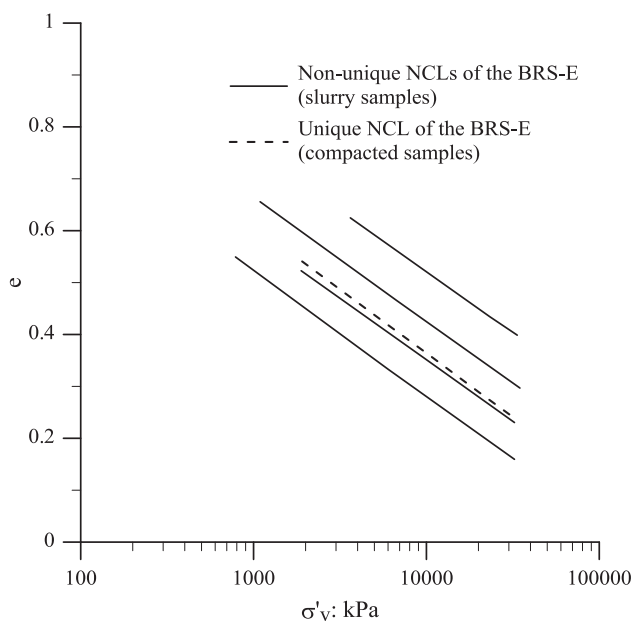


Fig. 13. Effect of sample preparation method on uniqueness of normal compression lines of sandy silt (E).

acted samples. It has a very similar slope to the compression lines of the slurry samples. Consequently, for the Bormida sandy silt (E), the sample preparation technique has an important effect on the presence of the transitional behaviour in one-dimensional compression.

#### 4. Conclusions

The one-dimensional compression behaviour of young alluvial soils, Bormida silty soils (BRS), has been investigated. Two different soil gradings were tested, namely, Bormida clayey silt (B) classified as ML and sandy silt (E) classified as CL-ML. The main important features of their behaviour are summarised in the following paragraphs:

- A unique normal compression line was found for the slurry samples of the clayey silt (B) regardless of the initial water content. This 1D-ICL\* was then used as a reference for analysing the effect of the initial structure of the compacted and the intact specimens.
- The normalized compression lines of the compacted samples of the clayey silt (B) were located to the right of the 1D-ICL\* and below the SCL, showing the effect of the initial structure created during compaction. The specimens with the highest initial void ratios reached the SCL, exhibiting a post-yield convergence towards the 1D-ICL\*. In the case of the densest samples, the effect of the initial structure was less pronounced. At high stress levels, and although the initial structure of each sample was expected to be different, the deformation induced during compression was enough to ensure that no differences in structure would be sufficient to significantly affect the meso-scale behaviour of the soil as they tended to converge onto the unique 1D-ICL\*. The NCL at high stress levels was very robust, regardless of the compaction method or the number and heterogeneity of the soil layers created during compaction and the presence of holes. This does not mean that the structure of each sample became the same for the range in stress applied.
- The effect of the in-situ structure of the intact block material was clear as the normalized compression curves crossed the 1D-ICL\* and yielded at different stress levels in between the 1D-ICL\* and the SCL. The stress sensitivity varied from 1.2 to 2.5 with the lowest and highest values corresponding to the densest and loosest samples, respectively. The post-yield behaviour showed a slow convergence towards the 1D-ICL\*, displaying a robust effect of the initial structure during one-dimensional compression, and at high stress levels, the stress sensitivity was still higher than unity. The absence of a clear and abrupt change in the slope of the compression lines after yield could be associated with there being only a small effect of the initial bonding on the soil behaviour. It appeared that the in-situ fabric of the Bormida clayey silt had a major effect on the compression behaviour.
- The intact and compacted effects of the structure were similar. This could lead to the conclusion that there will

be “no” structure effects at a meso-scale if the compacted behaviour is used as a reference. This confirms that the remoulded/compacted behaviour cannot be used as a reference for assessing the effect of the structure of the intact samples.

- The compression lines of the slurry specimens of the sandy silt (E) did not show any sign of convergence towards a unique NCL for the range in stresses applied. In fact, the lines ran parallel to each other and the locations appeared to depend on the initial structure of each sample. This behaviour means that the soil “remembers” its initial structure during compression even at large stresses, displaying clear transitional behaviour. This lack of uniqueness of the NCL would make it impossible to analyse the effect on the structure using the slurry behaviour as a reference. It also shows a lack of meaning for the term “intrinsic” behaviour as a reference framework. A unique NCL was identified for the remoulded compacted specimens regardless of the compaction method or the initial structure. The slope of the unique NCL defined by the compacted specimens was the same as those of the slurries. This result shows that, contrary to what is often stated in the literature, the sample preparation technique can affect whether a transitional mode of behaviour is seen or not.
- The fabric seemed to be the most important part of the structure in terms of controlling the behaviour of the young alluvial silts from the Bormida River. Similar behaviour has been seen in other young alluvial materials, such as Pisa clayey silt (e.g., Rampello and Callisto, 1998).
- The results have emphasised how the quantification of the natural structure is critically dependent on the way in which the reference samples are prepared.

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